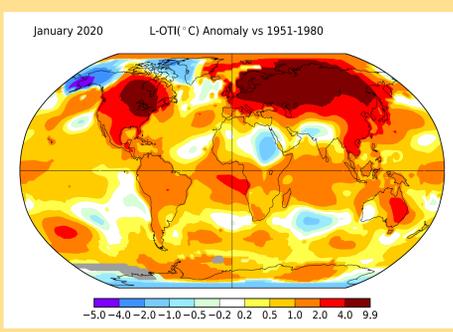


**Introduction**

- Arctic regions are warming much more rapidly than other parts of the globe<sup>1-2</sup> (Fig 1)
- Warming of this magnitude in this region is of concern due to the vast OC stores in permafrost<sup>3-4</sup> (frozen ground)
- Thaw ponds are the result of thawing permafrost peatlands and have been shown to emit CH<sub>4</sub> (Fig 2) yet there are few long-term studies of these water bodies<sup>5</sup>
- CH<sub>4</sub> from peatlands is emitted to the atmosphere via three pathways: (1) plant mediated transport, (2) hydrodynamic flux, and **ebullition (bubbling)**<sup>6-8</sup>
- Ebullition is often the dominant pathway but is also the least understood<sup>6-8</sup>



**Figure 1** Temperature in January 2020 compared to the mean January temperature during 1951 – 1980<sup>1-2</sup>.

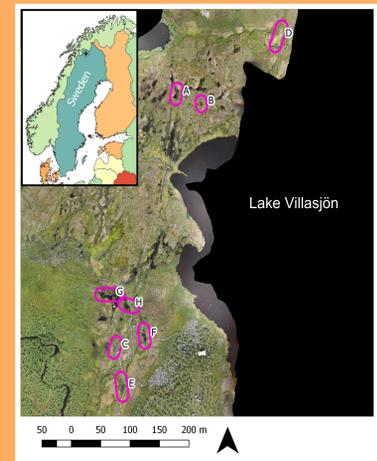
**Objective:** Investigating the seasonal and interannual controls of ebullitive flux from a subarctic thaw pond system



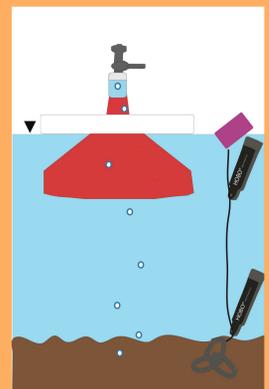
**Figure 2A-C** Development of Pond E through time. Photo 2A and 2B Credit: Patrick Crill.

**Methods**

- Eight ponds studied within Stordalen Mire, northern Sweden (68°21'N, 19°02'E) (Fig 3)
- Ponds sampled for bubbles during the growing season (June - September, every 1 – 3 days) in 2012 to 2015 with pond temperature measured continuously in all ponds since July 2013 (Fig 4)
- Samples analyzed on a Gas Chromatograph for CH<sub>4</sub> concentration (Fig 5)
- Drone imagery over Stordalen Mire from 2016 used to estimate pond size<sup>9</sup>
- Used non-parametric statistical tests to investigate drivers of ebullitive flux



**Figure 3** Pond sampling locations. Image Credit: Michael Palace, Jessica DelGreco, Christina Herrick.



**Figure 4** Schematic of funnel system and temperature loggers.



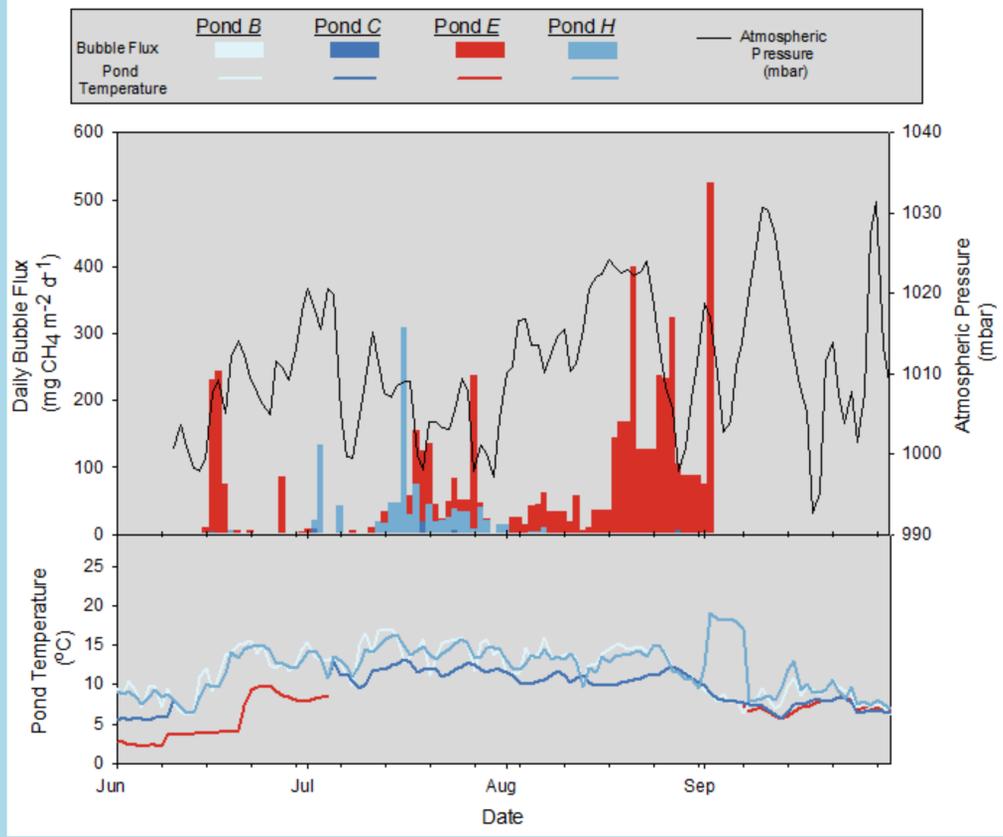
**Figure 5** Gas samples run on a Gas Chromatograph within 24 hours of collection. Photo credit: Clarice Perryman.

**Results**

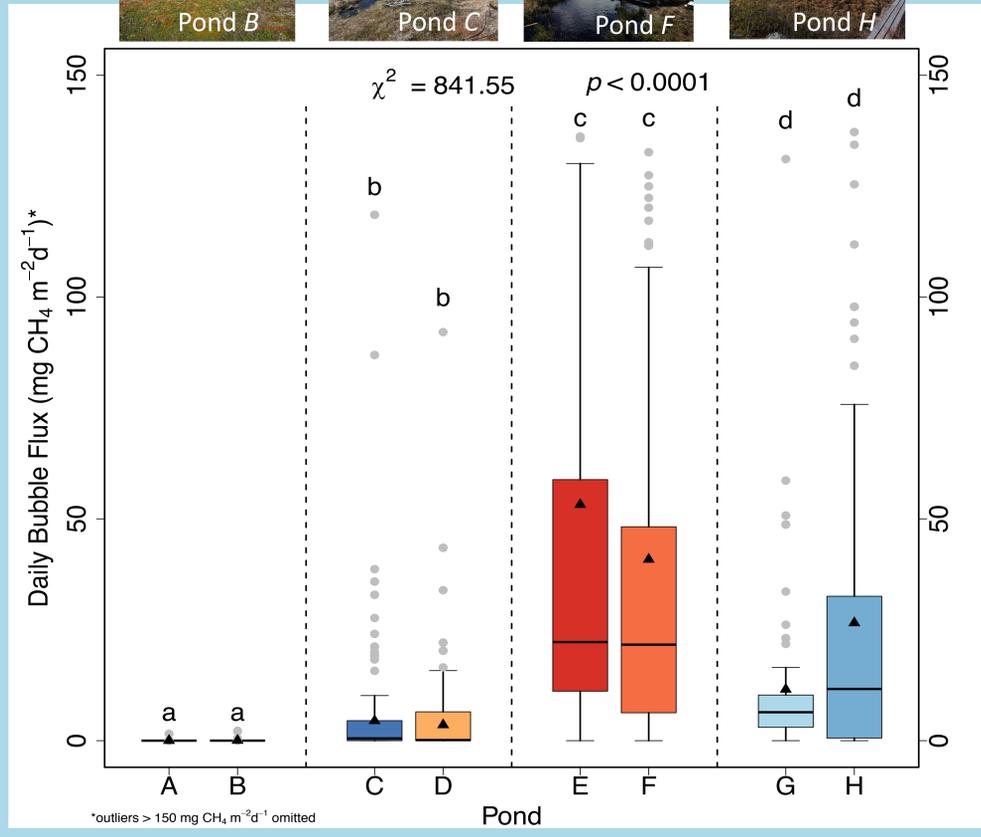
- Over the four-sampling season study, our eight ponds emitted on average **20 mg CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>**
- Episodic events (peaks in ebullitive flux) occurred in all ponds, sometimes associated with drops in atmospheric pressure but not always (Fig 6)
- Meteorological variables considered important to ebullition<sup>10-12</sup> correlated significantly with flux but did not explain much variability in flux
- Ebullitive flux varied significantly by pond, with the eight ponds appearing to fall into four distinct groups (Fig 7; Table 1)

**Table 1** Physical characteristics of the ponds broken down by type

	Type 1	Type 2	Type 3	Type 4
Depth (cm)	18 - 22	35 - 41	43 - 85	41 - 47
<i>Sphagnum spp.</i> present?	✓	✓	✓	✗
Sedges present?	<i>Eriophorum spp.</i>	↑ from Type 1	<i>Eriophorum spp.</i> & <i>Carex spp.</i> ↑ from Type 2	<i>Eriophorum spp.</i> & <i>Carex spp.</i> ↑ from Type 2
Hydrologic Connectivity <sup>13</sup>	Isolated	Transitioning	Isolated	Open Water; Connected to adjacent Fen



**Figure 6** Daily bubble flux (mg CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>) and pond temperature (°C) from selected ponds in 2015. Daily average atmospheric pressure (mbar) for each sampling season is displayed as a black line<sup>14</sup>.

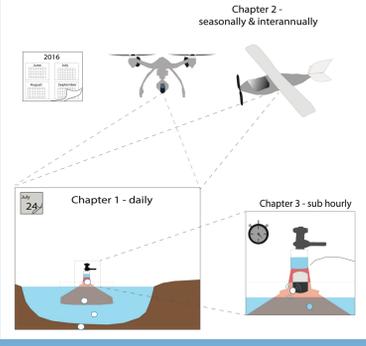


**Figure 7** Daily bubble flux 2012 – 2015 for each pond with mean daily bubble flux as ▲. Letters represent significant differences.  $\chi^2$  and  $p$  are from the Kruskal-Wallis rank sum test.

**Conclusions**

- Meteorological variables were very weakly correlated with thaw pond ebullition potential
- Using our average ebullitive flux, the estimated extent of small ponds above 50°N<sup>15-16</sup>, and an estimated 149 days ice-free season<sup>5</sup>, we estimate thaw ponds of < 0.001 km<sup>2</sup> to emit between 0.2 and 1.0 Tg CH<sub>4</sub>
- High frequency ebullition measurements over multiple seasons are important in constraining the CH<sub>4</sub> from small thaw ponds

**Figure 8** Schematic showing the various spatial and temporal scales of my PhD.



**Future Work**

Further examine ebullition from these ponds in both high temporal and spatial resolution using acoustic sensors and repeat drone fly overs (Fig 8).

**Acknowledgments** This work has been published in the AGU Journal of Geophysical Research: Biogeosciences and can be found here (<https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2018JG004786>). The coauthors would like to acknowledge the following funding in support of this project: the Northern Ecosystems Research for Undergraduates program (NSF REU site EAR 1063037, P. Varner), an NSF MacroSystems Biology Grant (EF 1241037, P. Varner), UNH's Hamel Center for Undergraduate Research SURF Abroad Grant, and by the Malmberg Scholarship Program of the American Swedish Institute, which is supported by the endowment established by Cornelia Malmberg through a generous bequest from her estate. A portion of this work was supported through the New Hampshire Space Grant Consortium, under NASA Grant NNX15AH79H. This study was in part funded by the Genomic Science Program of the United States Department of Energy Office of Biological and Environmental Research, Grants DE-SC0010580 and DE-SC0016440. This study has been made possible by the Swedish Infrastructure for Ecosystem Science (SITES), in this case at the Abisko Scientific Research Station (ANS). Meteorological data were provided by Swedish Polar Research Secretariat, ANS, Abisko, Sweden. Thanks to staff at ANS for providing housing, equipment, and laboratory space. Thanks to Niklas Rakos, Chris Hemmingsson, Matthew Osman, Justine Ramage, Mathilda Nyzell, Nathan Tomczyk, Eric Heim, Katharine Rocas, Kiley Remiszewski, Carmody McCalley, and James Lazarik for assistance in sample collection and analysis. Thanks to Grace Delgado for her knowledge and guidance with Adobe Illustrator. The figures in this paper are color-blind friendly using colors from www.ColorBrewer.org by Cynthia A. Brewer, Geography, Pennsylvania State University. Thanks to our reviewers for their diligent comments. The data associated with this work are stored on the data repository site of the ISOGENIE project and can be found online (<https://isogenie-db.asc.ohiostate.edu/datasources#fluxes>).

<sup>1</sup> GISTEMP Team, 2020. GISS Surface Temperature Analysis (GISTEMP). NASA Goddard Institute for Space Studies. Dataset accessed 2020-03-30 at <https://data.giss.nasa.gov/gistemp/>.  
<sup>2</sup> Hansen et al. (2010). Rev. Geophys. <sup>3</sup> Loisel et al. (2014). The Holocene <sup>4</sup> Yu et al. (2010). Geophys. Res. Lett. <sup>5</sup> Wik et al. (2016). Nat. Geosci. <sup>6</sup> Bestviken et al. (2011). Science. <sup>7</sup> Coulthard et al. (2009). In Carbon Cycling in Northern Peatlands (pp. 173-186). <sup>8</sup> Fechner-Levy & Hemond (1996). Limnol. Oceanogr. <sup>9</sup> Palace et al. (2018). Remote Sens. <sup>10</sup> Goodrich et al. (2011). Geophys. Res. Lett. <sup>11</sup> Weyhenmeyer (1999). Global Biogeochem. Cycles. <sup>12</sup> Wik et al. (2014). Geophys. Res. Lett. <sup>13</sup> Orlfeldt & Roulet (2012). J. Geophys. Res. <sup>14</sup> ANS. (2017). Abisko Weather Station Data. [A. S. R. Station, Ed.]. Abisko Scientific Research Station. Retrieved from <http://gaisler.se/abisko>. <sup>15</sup> Holgersson & Raymond (2016). Nat. Geosci. <sup>16</sup> Verpoorter et al. (2014). Geophys. Res. Lett. <sup>17</sup> Burke, S. A., Wik, M., Lang, A., Contosta, A. R., Palace, M., Crill, P. M., & Varner, R. K. (2019). Long-Term Measurements of Methane Ebullition From Thaw Ponds. Journal of Geophysical Research: Biogeosciences, 2018JG004786. <https://doi.org/10.1029/2018JG004786>