

Drivers of Temporal Variability in Ocean Iron Stress in HNLC Regions using Ocean Color and Climate Modes



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Motivation

- Phytoplankton are microscopic plants that carry out photosynthesis, produce 50-80% of Earth's oxygen, and is food for marine animals.
 - Photosynthesis requires micronutrient iron
 - Over 1/3 of the global ocean is iron limited (HNLC)
 - Ocean warming is predicted to limit iron supply
 - Impact oxygen supply & marine resources

13 Monitor long-term climate change impacts on phytoplankton

14 Conserve marine resources for future ocean health

- Ship-based time-series measurements of dissolved Fe from high nutrient low chlorophyll (HNLC) regions (Equatorial Pacific, Subarctic N. Pacific, and Southern Ocean) are limited to nonexistent.
- Objective: To determine the temporal variability (interannual & seasonal) of NPQ-corrected ϕ_{sat} (Eq 1), a satellite remote sensing-based estimate of phytoplankton iron (Fe) stress, across HNLC regions with related climate modes and other satellite remote-sensed products.**

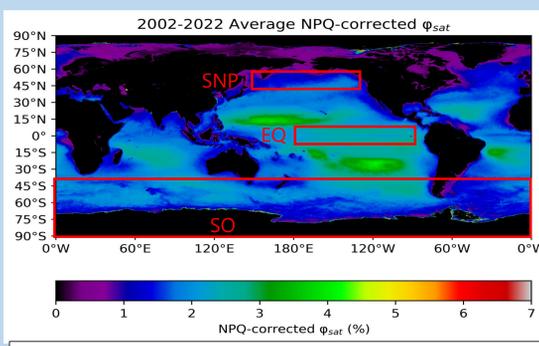


Figure 1. Global 20-yr average NPQ-corrected ϕ_{sat} . HNLC regions (red boxes) are nutrient replete, Fe limited, sunlit waters with relatively low rates of biological production/export. Subtropical gyres have chl < 0.1 or other nutrient stresses.

Background

- Phytoplankton pigments absorb light for photosynthesis, but excess light is fluoresced and emitted as red (675 nm).
 - Fluorescence is also an indicator of chlorophyll (chl) concentration, photosynthetic rates, and physiological status^{1,2,7} (Eq 2).
- High fluorescence yields are associated with physiological and nutrient stress in phytoplankton⁸.
- By correcting for (1) pigment concentration, (2) photoprotective response or nonphotochemical quenching (NPQ, heat dissipation), and (3) pigment packaging, **NPQ-corrected ϕ_f can be derived as a proxy for Fe stress**^{1,7}.
- Under non-macronutrient-stressed conditions (N and P abundance), iron stress in chl increases fluorescence quantum yields (ϕ_f)^{1,5}.
- Satellite-derived ϕ_{sat} can indicate phytoplankton Fe stress¹.

Equations

Eq 1: NPQ-corrected $\phi_{sat} = 0.00043 \frac{nFLH \cdot iPAR}{Chl^{0.684}} * 100$

Eq 2: Fluorescence (F) = $iPAR * Chl * \alpha_{phyto}^* \phi_f$

- iPAR**: Instantaneous photosynthetically available radiation
- α_{phyto}^*** : chl-specific phytoplankton absorption coefficient
- ϕ_f (phi)**: Quantum yield of fluorescence
 - Ratio of light fluoresced by chlor-a pigments to total light absorbed by phytoplankton

Eq 3: Particulate Carbon (PC) = PIC + POC

Eq 4: Coefficient of variation (CV) = $\frac{std\ dev}{average} * 100$

Eq 5: SST anomaly = $SST_{month} - SST_{monthly\ 20-yr\ average}$

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Results and Discussion

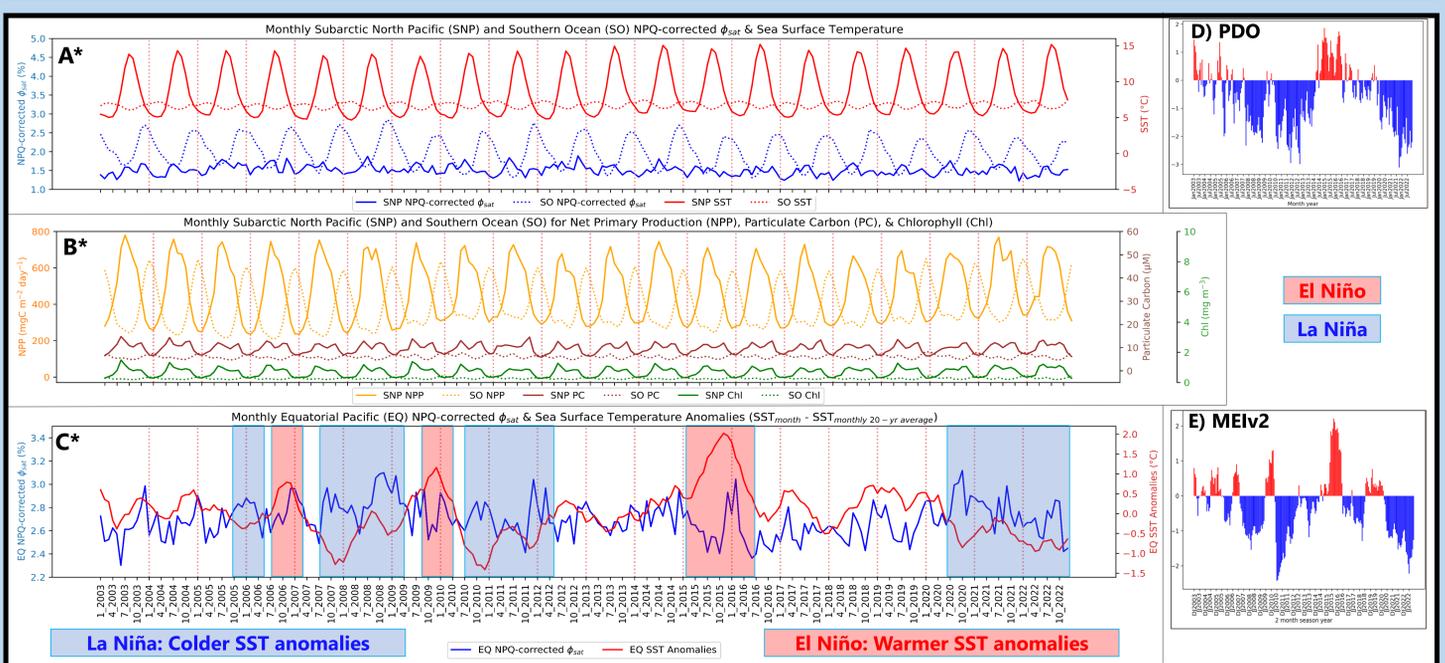


Figure 2. Monthly (A) NPQ-corrected ϕ_{sat} and SST by SNP (solid line) and SO (dotted line); (B) NPP, PC, and chl by SNP and SO; (C) EQ NPQ-corrected ϕ_{sat} and SST anomaly; (D) PDO; (E) MEIv2 climate modes for 2002-2023. *Vertical dotted lines are one-year increments.

- No secular trends in Fe stress (Blue lines). SNP and SO are subject to seasonal cycles. SNP sea ice impact during November-March with a freshwater cap/stratification that prevents convective mixing in the winter¹⁰. East SNP gets more Fe inputs than west SNP¹⁰. SO sea ice impact during April-October.
- EQ Fe stress is tied to El Niño-Southern Oscillation (ENSO); greater Fe stress with La Niña events (colder SST anomalies); El Niño 2015 (exception) had a lagging elevated Fe stress response with warmer SST anomalies.

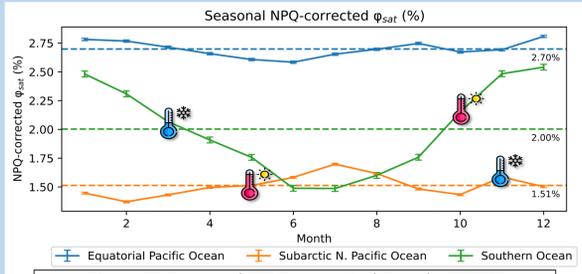


Figure 3. Seasonal NPQ-corrected ϕ_{sat} by HNLC. Winter starts. Summer starts.

Table 1. Seasonal statistics.

	EQ	SNP	SO
Ave ϕ_{sat} (%)	2.70	1.51	2.00
Std Dev	0.16	0.13	0.39
Std Err	0.01	0.01	0.03
CV (%)	5.76	8.51	19.56

Fe stress: EQ>SO>SNP
 • SNP peaks in July (summer)
 • SO peaks in Dec (summer)
 Fe stress variability: SO>SNP>EQ
 • SO: highest variability (19.56%)
 • EQ: highest stability (5.76%)

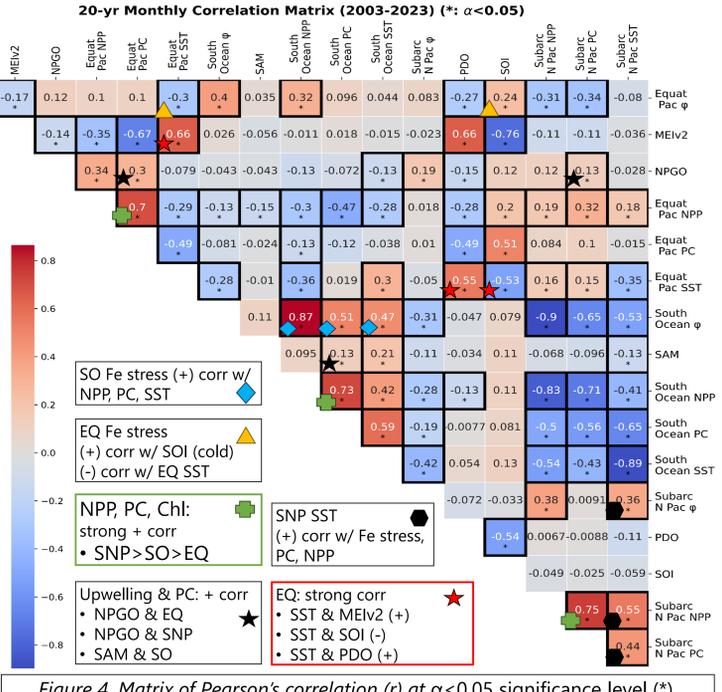


Figure 4. Matrix of Pearson's correlation (r) at $\alpha < 0.05$ significance level (*).

Methodology and Analysis

HNLC regions

- SNP**: Subarctic North Pacific: 60°N-40°N; 140°E-120°W
- EQ**: Equatorial Pacific Ocean: 10°N-10°S; 180°E-85°W
- SO**: Southern Ocean: 40°S-90°S

Monthly, Seasonal, Interannual Analysis

- Monthly averages by HNLC region (Fig. 2)
- Monthly 20-yr averages for HNLC regions (Fig. 3)
- Seasonal statistics (Table 1)
- Linear correlation coefficients with Pearson's correlation (r) at $\alpha < 0.05$ significance level (Fig. 4)

Climate Modes

- PDO**: Pacific Decadal Oscillation: ~SNP
 - SST EOF
 - +: El Niño & -: La Niña
- NPGO**: North Pacific Gyre Oscillation: ~SNP
 - SSH EOF
 - +: Upwelling & -: Downwelling
- SOI**: Southern Oscillation Index: ~EQ
 - SLP difference btw Tahiti and Darwin
 - +: La Niña & -: El Niño
- MEIv2**: Multivariate ENSO Index: ~EQ
 - SST, SLP, wind (U & V), outgoing thermal radiation EOF
 - +: El Niño & -: La Niña
- SAM**: Southern Annular Mode: ~SO
 - Zonal sea level pressure (SLP) means at 40°S and 65°S
 - +: Upwelling & -: Downwelling

MODIS-Aqua Products

Two decades (2003-2022) of Moderate-resolution Imaging Spectroradiometer (MODIS)-Aqua mapped monthly Level-3 products at 9 km spatial resolution from NASA ocean color (<http://oceancolor.gsfc.nasa.gov>)

- Chl_{sat}**: mg chl m⁻³ (Detection limit < 0.1)⁴
- iPAR_{sat}**: $\mu\text{mol photons m}^{-2} \text{s}^{-1}$
- nFLH_{sat}**: mW cm⁻² $\mu\text{m}^{-1} \text{sr}^{-1}$
 - (Detection limit < 0.003)⁶
 - normalized fluorescence line height
- PIC_{sat}**: particulate inorganic carbon (μM)
- POC_{sat}**: particulate organic carbon (μM)
- SST_{sat}**: sea surface temperature (°C)
- NPP_{sat}**: net primary productivity ($\text{mg C m}^{-2} \text{day}^{-1}$)
- CAFÉ algorithm⁹.

- In the EQ, Fe stress decreases as SST increases (positive phase MEIv2 & PDO; negative phase SOI)
 - Fe stress driven by La Niña periods (colder SST) (Fig. 2C).
 - No correlation among NPP and PC with Fe stress (Fig. 4).
- In the SNP and SO, as SST increase, NPP & PC increases, which increases Fe stress as nutrients are taken up (Fig. 2A).
 - Seasonal wintertime reduced Fe stress.
- As upwelling increases (+NPGO) in SNP, stimulated phytoplankton blooms contribute to higher dissolved Fe uptake and increased Fe stress following Fe depletion.
 - Previous research for the EQ found upwelling supplies Fe, thus, reducing Fe stress and increasing chlorophyll concentrations³.
 - Our analysis expands this: **following nutrient input (upwelling) and/or seasonal warming, Fe stress may temporarily decrease, but it increases at carrying capacity.**
- Higher stability and minimal seasonal variability in the EQ may select phytoplankton with lower Fe requirements, while polar regions have greater seasonal shifts that cause restructuring of the phytoplankton community^{3,5}.

Conclusions

- Fe stress varies interannually based on supply, use, & SST.
- Seasonal summer cycle (\uparrow SST, \uparrow NPP) enhances Fe stress in high latitudes at carrying capacity with respect to Fe supply.
 - Upwelling phase correlated with Fe stress in SNP.
- Lowest Fe stress in SNP, but highest PC & NPP, indicates more Fe input supply than EQ & SO.
- Highest Fe stress variability in SO due to sea ice cover
- Cooler SST/ La Niña cycle drives Fe stress in EQ
- EQ has the highest Fe stress and high NPP, but lowest PC, indicating faster carbon cycling.
- Not all variation can be explained by climate modes
 - i.e. phytoplankton taxonomic composition, accessory pigmentation, physical circulation, and Fe inputs¹⁰.

Future Work

- Focus on localized limitation in narrower latitudinal bands.
- Conduct validation tests of NPQ-corrected ϕ_{sat} with cruise data.
- Tease out changes in phytoplankton community composition as adaptation and community shifts occur due to ocean warming and environmental responses with NASA's satellite PACE (Plankton, Aerosol, Cloud, ocean Ecosystem).